

# Spatial Variation of Textural Parameters in a Small River: An Example from Khurar River, Khajuraho, Chhaterpur District, Madhya Pradesh, India

S. Kanhaiya and B.P. Singh\*

*Centre of Advanced Study in Geology, Banaras Hindu University, Varanasi-221005, India*

**Abstract:** Texture (grain size) is one of the many parameters, which is used in determining depositional environments of sediments and sedimentary rocks. In the present investigation, texture was analysed from sixteen stations in the Khurar River, Madhya Pradesh, India. Here, grain size analysis was carried out employing mechanical sieving method using a sonic shaker. Frequency and cumulative frequency curves were prepared from the grain size data on centimetre and arithmetic probability papers, respectively. The phi values were determined and used to calculate the statistical parameters such as mean, standard deviation, skewness and kurtosis. It is found that the mean size value varies from 0.63 to 0.80 with a graphic mean distribution ranging from -0.27 to 0.40  $\phi$ , indicating that the size of the river sand is very coarse to coarse-grained. The standard deviation (sorting) shows a range of 0.69 to 1.65  $\phi$ . The skewness values of the sediment samples range from 0.19 to 0.29  $\phi$ , thus, indicating the presence of fine fraction to near-symmetrical fraction in the population. The kurtosis varies between 1.03 and 1.09  $\phi$ , indicating that 25% of the samples are leptokurtic, 6.25% are very leptokurtic, 50% are mesokurtic and 12.5% are platykurtic. The platykurtic nature in few cases suggests mixing of sediments from two sources. Bivariate plots prepared combining various textural parameters were used to interpret their behaviour in the river sediments. C-M plot was also prepared to understand the dominant mode of sediments transportation in the Khurar River. In this river, all the sediments are dominantly characterized by the rolling process of transportation. This study reveals that sorting varies from poorly sorted to moderately well sorted in the course of the river may be because of dominance of winnowing and selective sorting in the lower reaches of the river.

**Keywords:** Grain size, Kurtosis, Standard deviation, Sieving, Skewness, Sonic shaker.

## 1. INTRODUCTION

Texture (grain size) is the most fundamental property of sediment particles, affecting their entrainment, transport and deposition [1]. It helps in determining depositional environment, besides others. Grain size analysis provides important clues to the sediment transportation history, depositional conditions and provenance [2-4]. The downstream distribution of river bed sediments is often used to reflect regional fluvial geomorphological and hydrological settings [5]. There are various techniques for the determination of grain size, such as direct measurement, dry and wet sieving, sedimentation, and measurement by laser granulometer, X-ray sedigraph and Coulter counter. In this study, the textural analysis was carried out involving sieving technique. Stream sediments of Khurar river have been collected (starting from up-stream to down-stream direction) and analysed. Grouped data of grain sizes, obtained by sieving technique, are presented in a Table 1. The attributes of grain-size frequency distribution are then computed graphically following [2, 3]. Different modes of sediment transportation have been identified by plotting the cumulative frequency distributions of grain sizes on

probability paper. The net distribution is then divided into traction load (coarsest fraction), saltation load (medium fraction) and suspension load (finest fraction).

## 2. GENERAL AND GEOLOGICAL FEATURES OF THE RIVER

The Khurar River flows exclusively on Archean rocks of Bundelkhand carton. This comprises over ~29,000 km<sup>2</sup> area in north-central India. The carton consists of slivers of Archean greenstone successions within granitoids and gneisses, latter containing components with ages 3.5 Ga [6], 3.3 Ga, 2.7 Ga and 2.56-2.49 Ga [7]. The Khurar River originates from the village Saddapura (Maharajanj) near Khajuraho town in Chhaterpur district, Madhya Pradesh, India. This river is one of the tributaries of Ken River that is a tributary of river Yamuna. It is a small river and the length of the river is about 34 kilometre, and originates from the Beni sagar reservoir (Lat. 24°48' 3.5" N and Long. 79°52'59.6" E) and confluences with Ken river at Ghadiyal pond, near renneh fall in the area of Panna tiger reserve (Lat. 24°54' 13.9" N and Long. 80°2'6" E) and flows from SW to NE direction (Figure 1). The width of the river is ranging from 5 to 15 meter and average velocity is 2.5 meter/second. The river deposits small (2-4m. large) braid-bars in the middle of the channel that deposits sediments during monsoon season mainly.

\*Address corresponding to this author at the Centre of Advanced Study in Geology, Banaras Hindu University, Varanasi-221005, India;  
Tel: 915426701355; Fax: 915422369239;  
E-mail: drbpsingh1960@gmail.com

Table 1: Percentile Values, C-M Values and GRAPHIC MEASURES calculated from the Grain-Size Data

Sample Nos.	Φ95	Φ84	Φ75	Φ50	Φ25	Φ16	Φ5	Φ1	C in micron	M in micron	M <sub>z</sub>	σ <sub>1</sub>	SK	KG	σs	SoS
K <sub>1</sub>	2.8	1.5	0.7	-0.6	-1.4	-1.7	-2.5	-3.4	10500	1510	-0.27	1.65	0.29	1.03	1.50	0.09
K <sub>2</sub>	2.5	1.3	0.7	-0.8	-1.9	-1.9	-3.1	-4.2	18300	1740	-0.47	1.65	0.24	0.96	1.00	0.09
K <sub>3</sub>	2.1	1.5	1.2	0.4	-0.1	-0.4	-1.8	-3.2	9100	750	0.50	1.06	0.02	1.23	-0.50	0.13
K <sub>4</sub>	0.7	0	-0.3	-0.4	-0.9	-1.1	-1.9	-2.1	4200	1320	-0.50	0.87	-0.23	1.77	0.40	0.19
K <sub>5</sub>	1	0.4	0	-0.5	-1.2	-1.5	-2.1	-2.9	7400	1410	-0.53	0.94	-0.04	1.06	-0.10	0.16
K <sub>6</sub>	1.2	0.5	0.2	-0.4	-1.1	-1.4	-2	-2.7	6400	1320	-0.43	0.95	-0.02	0.69	0.00	0.15
K <sub>7</sub>	0.9	0.3	0	-0.6	-1.2	-1.4	-2	-2.7	6400	1510	-0.56	0.86	0.05	0.99	0.10	0.17
K <sub>8</sub>	1.8	1	0.7	0.2	-0.3	-0.4	-1	-1.7	3200	870	0.60	0.77	0.14	1.14	1.20	0.18
K <sub>9</sub>	1.9	0.9	0.6	-0.2	-0.9	-1.2	-2	-3.0	8000	1150	-0.16	0.84	0.04	1.06	0.30	0.12
K <sub>10</sub>	1.3	0.8	0.4	0.1	-0.4	-0.7	-1.6	-2.7	6400	930	0.03	0.81	-0.11	1.48	-0.50	0.17
K <sub>11</sub>	1.2	0.8	0.7	0.2	-0.5	-0.8	-1.5	-2.4	5200	860	0.06	0.81	-0.25	0.92	-0.70	0.18
K <sub>12</sub>	1.4	0.9	0.6	0.1	-0.4	-0.6	-1.6	-2.6	6060	930	0.13	0.83	0.03	1.22	-0.40	0.17
K <sub>13</sub>	0.8	0.1	-0.1	-0.6	-1.2	-1.4	-2	-2.6	6060	1510	-0.63	0.79	-0.03	1.04	0.00	0.18
K <sub>14</sub>	1.5	0.9	0.6	0.1	-0.4	-0.9	-1.9	-3.1	8500	930	0.03	0.96	-0.14	1.39	-0.60	0.15
K <sub>15</sub>	2.3	1.7	1.5	0.9	0.1	-0.2	-1.1	-1.9	3700	530	0.80	0.98	-0.15	0.87	-0.60	0.14
K <sub>16</sub>	0.9	0.9	0	-0.5	-0.9	-1	-1.5	-2.0	4000	1410	0.40	0.69	0.198	1.09	0.40	0.21

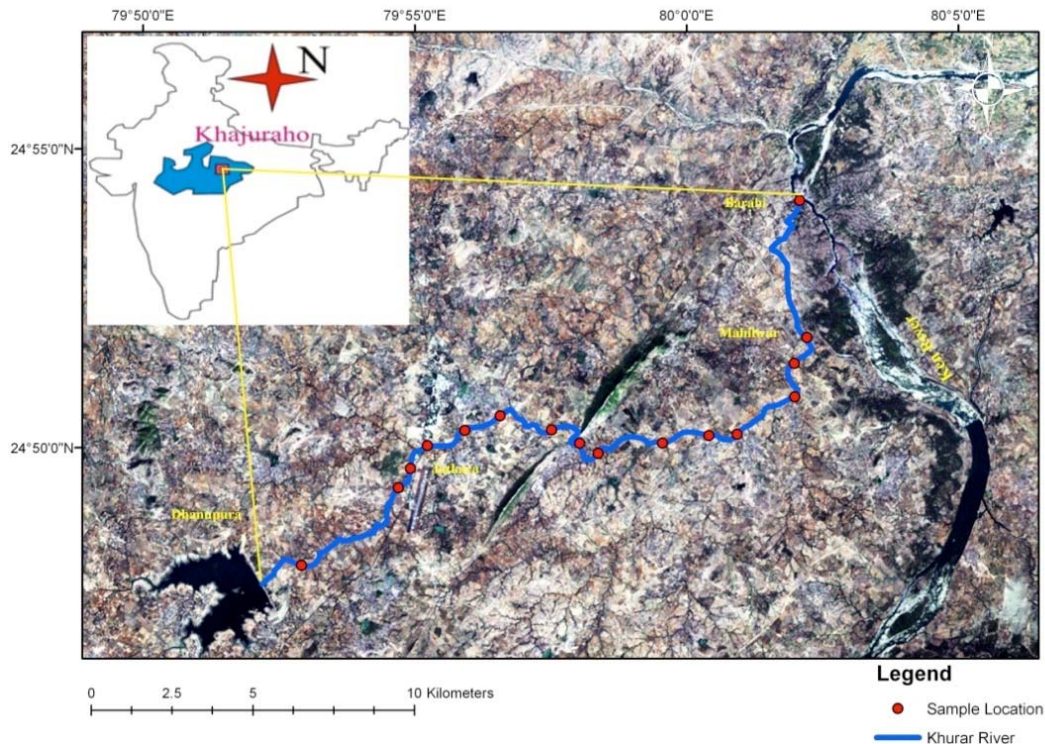


Figure 1: Satellite imagery showing the course of the Khurar River and sampling locations in the course of the river.

3. SAMPLING AND METHODOLOGY

A total of forty-eight sediment samples from sixteen stations those occurs on the braid-bars were collected in the entire course of the Khurar River at an interval of

about one kilometre (Figure 1). Sampling sites were positioned by GPS and all the samples were collected from the middle part of the channel. About two hundred gram by weight of the sample was collected in each case. One each sample from each station was utilized

for sieving analysis by coning and quartering method. Sieving was carried out at half phi ( $\phi$ ) intervals starting from 10 mesh (-1  $\phi$ ), 14 mesh (-0.5  $\phi$ ), 18 mesh (0  $\phi$ ), 25 mesh (0.5  $\phi$ ), 35 mesh (1.0  $\phi$ ), 60 mesh (2.0  $\phi$ ), 120 mesh (3.0  $\phi$ ), 170 mesh (3.5  $\phi$ ), 230 mesh (4.0  $\phi$ ), and shook for fifteen minutes followed by weighing of each retained fraction. Individual weight percentages were recalculated to hundred for making them in percent. The frequency curves were prepared on simple graph paper. Cumulative curves were plotted on arithmetic probability graph paper as proposed by [8-12]. The graphic parameters proposed by [2], C-M plot to evaluate the hydrodynamic force working during the deposition proposed by [13] and log-probability values proposed by [3] were calculated from the percentile values (in  $\phi$  units) of the cumulative curves. Bivariate plots were plotted considering various textural parameters along with spatial variation of sorting.

#### 4. GRAIN SIZE DISTRIBUTION

Grain size analysis is a classical tool and provides additional information regarding sediment transport, energy conditions and depositional environment [14-19]. Various textural parameters such as graphic mean, standard deviation (sorting), skewness and kurtosis have environmental significance and are useful for understanding synsedimentary hydrodynamic factors of transportation and deposition in a basin [2, 3, 19-22].

#### 4.1. Cumulative Frequency Curves

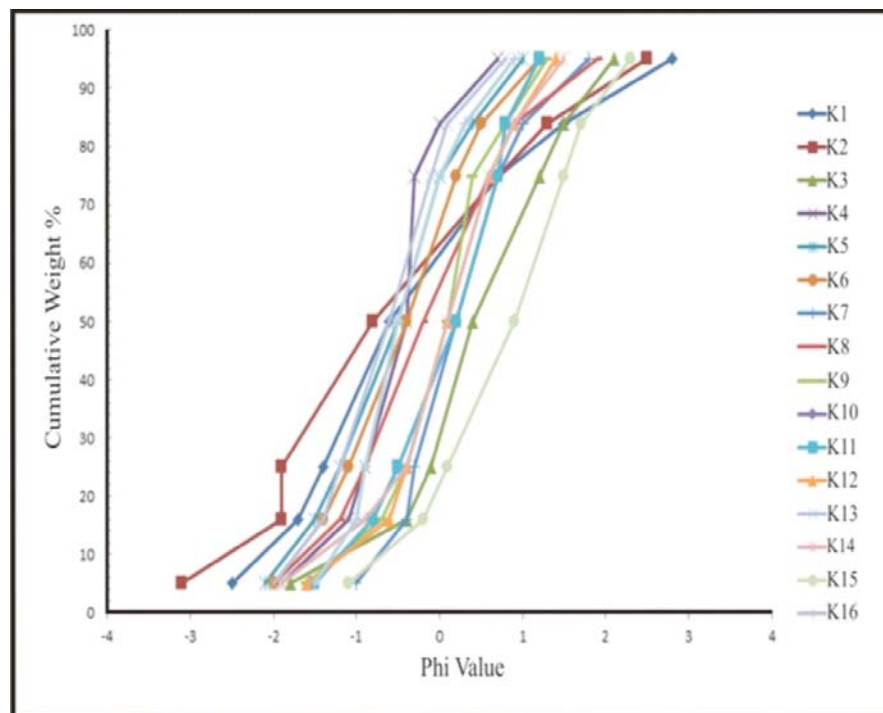
Frequency curves show that most of the samples are bimodal in nature, except K4, K8, K15, and K16 (Figure 2). Also, they show that the sediment population are of very coarse to coarse-grained. The cumulative frequency curves show three to four segments in each case representing traction, saltation I, saltation II and suspension fraction within the sediments. In all, traction and saltation load dominates over suspension load.

#### 5. TEXTURAL PARAMETERS

All the textural parameters were calculated from cumulative curves, based on different phi values. A number of formulas have been proposed by different workers to calculate four main statistical parameters viz. graphic mean, graphic standard deviation, graphic skewness and graphic kurtosis [15, 20, 23-25]. Formulas given by Folk and Ward (1957) are regarded as most suitable and used for calculation of different textural parameter in this paper (Table 1).

##### 5.1. Graphic Median

Graphic median values i.e.  $\phi_{50}$  of all individual samples have been analysed statistically, denoting that half of the particles by weight are very coarse to it and half are coarse grained. The obtained values range



**Figure 2:** Cumulative curves showing the trends of all the sediment samples.

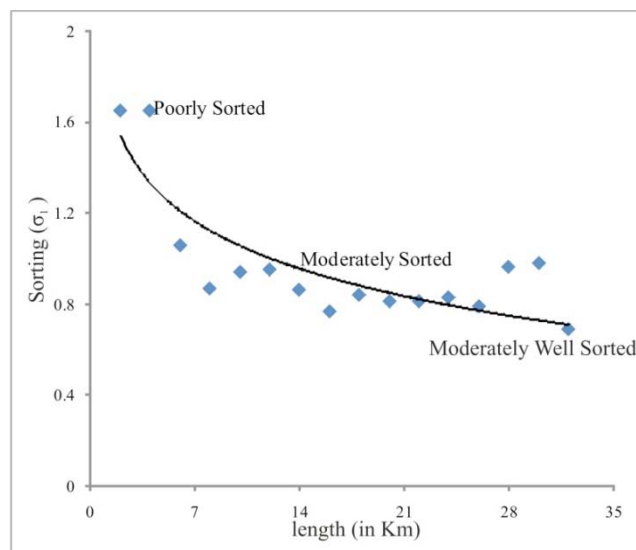
from  $-0.08\phi$  to  $0.90\phi$ , averaging  $-0.16\phi$ . The average value of median in individual sediment samples, shows the dominance of very coarse-grained sediments. However, out of sixteen, seven samples i.e. K3, K8, K10, K11, K12, K14 and K15 show large percentage of coarse-grained sediments with  $\phi_{50}$  values 0.4, 0.2, 0.1, 0.2, 0.1, 0.1, and  $0.9\phi$  respectively.

## 5.2. Graphic Mean

Graphic mean ( $M_z$ ) is a measure of an arithmetic average of a series of values and is calculated by the formula  $\phi_{16} + \phi_{50} + \phi_{84} / 3$ . The calculated values range from  $-0.63$  to  $0.80\phi$  with an average value of  $-1.1\phi$ . The average value of mean of all individual samples denotes that the major sediment class is very coarse grained. But few individual samples i.e. K3, K8, K10, K11, K12, K14, K15 and K16 have comparatively large fractions of coarse-grained sediments (Figure 4a).

## 5.3. Graphic Standard Deviations

The graphic standard deviation ( $\sigma_1$ ) is a measure of sorting or variation in sizes and is calculated by the formula  $\phi_{84} - \phi_{16} / 4 + \phi_{95} - \phi_5 / 6.6$ . The values obtained range from 0.69 to  $1.65\phi$ . Samples are poorly sorted near the source of the river following a trend of moderately shorted as we go towards the downstream side and moderately well sorted near to the confluence of the river with Ken River (Figure 3). Out of sixteen, some samples such as K1, K2 and K3 are poorly sorted having  $\sigma_1$  values 1.65, 1.65, and  $0.50\phi$ , but K4, K5, K6, K7, K8, K9, K10, K11, K12, K13, K14 and K15



**Figure 3:** Spatial variation of the sorting values along the length of the Khurar River.

are moderately sorted having  $\sigma_1$  values ranging from 0.77 to  $0.98\phi$  whereas last sample near the point of confluence with Ken River K16 is moderately well sorted having  $\sigma_1$  value  $0.68\phi$  (Figure 4b).

## 5.4. Graphic Skewness

The graphic skewness ( $Sk_1$ ) deals with the quality, state, or condition of being distorted or lacking symmetry. It is the measurement of systematic distribution or predominance of coarse or fine sediments and usually calculated by the formula  $\phi_{84} + \phi_{16} - 2\phi_{50} / (\phi_{84} - \phi_{16}) + \phi_{95} + \phi_5 - 2\phi_{50} / (\phi_{95} - 2\phi_5)$ . The skewness values range from  $-0.23\phi$  to  $+0.29\phi$  i.e. near-symmetrical to very fine-skewed. However, there is a dominance of near-symmetrical category in the samples of Khurar River. Seven samples i.e. K3, K5, K6, K7, K9, K12 and K13 are near-symmetrical followed by five samples i.e. K4, K10, K11, K14 and K15 are coarse-skewed and four samples i.e. K1, K2, K8 and K16 are of very fine-skewed (Figure 4c).

## 5.5. Graphic Kurtosis

The graphic kurtosis (KG) is defined as the peakedness of the distribution and measures the ratio between the sorting in the tails and central portion of the curve. It is calculated by the formula  $\phi_{95} - \phi_5 / 2.44(\phi_{75} - \phi_{25})$ . If the tails are better sorted than the central portions, then it is termed as platykurtic, whereas, leptokurtic, if the central portion is better sorted. If both are equally sorted then mesokurtic condition prevails.

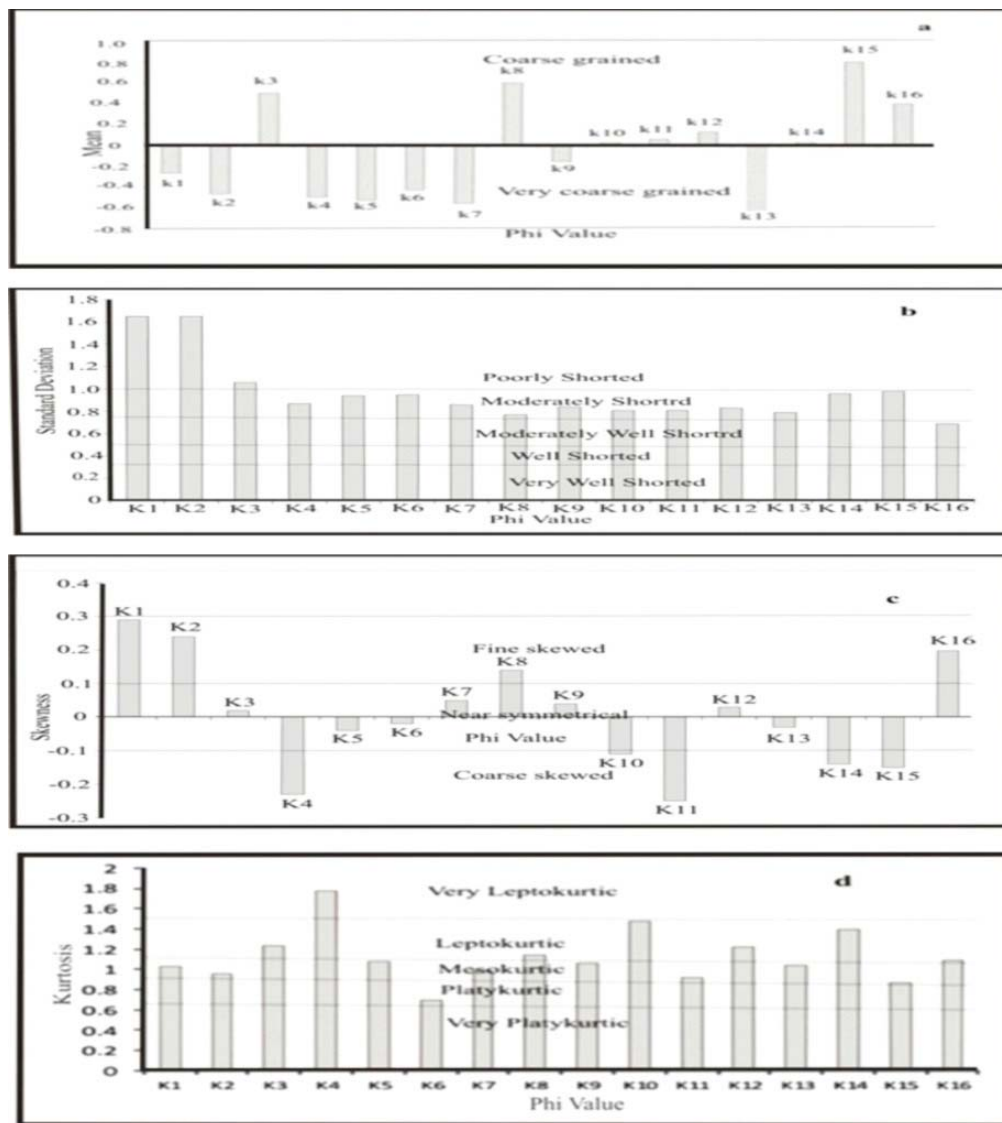
The values obtained range from 0.69 to  $1.77\phi$ , however, there is a dominance of mesokurtic (a total of eight samples, values ranging from 0.96 to  $1.09\phi$ ) to leptokurtic (a total of seven samples, values ranging from 1.14 to  $1.77\phi$ ). Only two samples K6 ( $0.69\phi$ ) and K15 ( $0.87\phi$ ) are platykurtic in nature (Figure 4d).

## 6. INTER-RELATIONSHIP OF TEXTURAL PARAMETERS

The inter-relationship between different parameters is recognisance tool to interpret various aspects of depositional environment, as the textural parameters of the sediments are often environmentally sensitive [2, 20, 21, 24, 26, 27]. Following Folk and ward [2], six bivariate scatter plots are prepared by the combination of various textural parameters.

### 6.1. Mean versus Standard Deviation

The Mean versus standard deviation plot gives a great amount of information about the depositional



**Figure 4:** Comparative histograms of all samples showing variations of (a) mean, (b) standard deviation, (c) skewness and (d) kurtosis values.

environment [1, 19, 27]. If a wide range of grain size (gravel to clay) is present scatter band often form some segment of a broadened M-shaped trend, only a V-shaped or inverted V-shaped trend develops if the size range is smaller, and if range is very small, only one limb of V may occur [2]. The samples of study area shows inverted V-shape trend due to small particle size range. The distribution pattern shows the clustering of values slightly deviated from the middle position, tending towards the right limb of inverted V-shaped established trend of [2], indicating a smaller size range of the grains (Figure 5a). The attention of sediment at right limb of V-shape plot suggests mixture of the dominant sand mode with a small amount of silt. This may also be reflected dominantly bimodal and moderately-sorted nature of sediments.

## 6.2. Mean versus Skewness

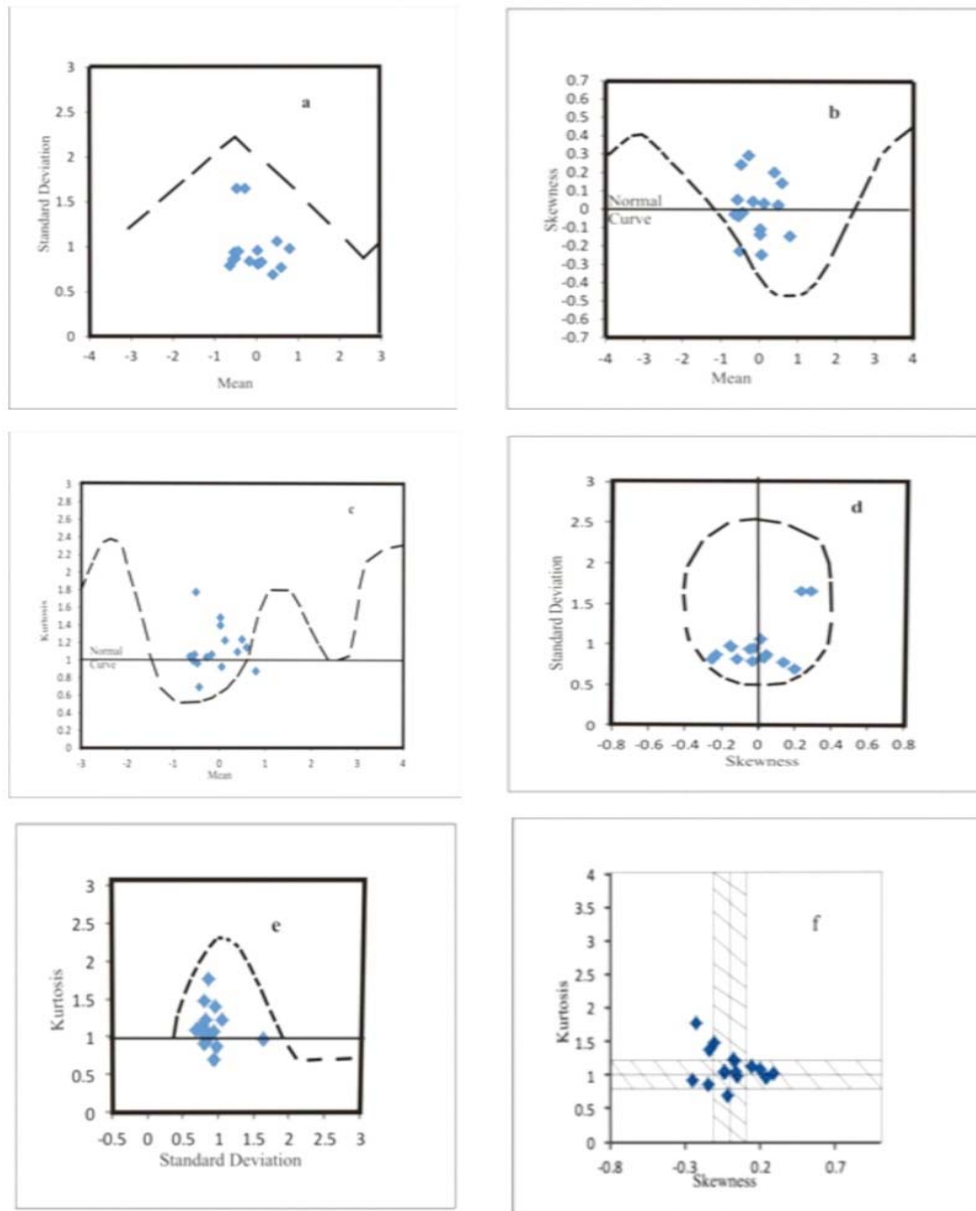
The sinusoidal curve of mean versus skewness following [2] was plotted. The sinusoidal nature is because of proportionate admixture of two size-classes of the sediments. In general, the ideal fractions are nearly symmetrical but the mixing produces either positive or negative skewness depending upon the proportions of size-classes in the admixture. According to Friedman [20], river sand is generally positively skewed whereas coarse grained river sand (coarse sand to gravel) may either be positively skewed or negatively skewed [28,29]. In this study the near-symmetrical values are almost confined equally in the positive-skewed area of the graph, as well as negatively skewed, in the mean-size range of  $-1$  to  $1\phi$  (Figure 5b). It strongly suggests a bimodal nature of

sediments and mixture of the dominant sand mode with a very small amount of silt.

**6.3. Mean versus Kurtosis**

The relation between mean-size and kurtosis is complex and theoretical [2]. The model plot of Folk and Ward [2] denotes the mixing of two or more size-classes of sediments, which basically affects the sorting in peak and tails i.e. index of kurtosis. In the present study scattering pattern of all the samples gives rise to a V-shape trend. The nature of the plot also indicates a dominance of mesokurtic (0.90 to

1.11 $\phi$ ) nature of sediments followed by platykurtic (0.90 to 0.67 $\phi$ ) and leptokurtic (1.11 to 1.50  $\phi$ ). The size of all the particles is showing a range of approximately -1 to 1 $\phi$  i.e. very coarse to coarse sand (Figure 5c). It also confirms that the sediment-admixture is dominated by sand (coarse to very coarse) and small proportion of silt. The mixing of sediments with dominant sand mode and some proportion of fine sediment makes the sorting worse, particularly in the tails; hence, there is a presence of mesokurtic along with platykurtic and leptokurtic conditions.



**Figure 5:** Bivariate plots of the sediment samples of the Khurar River (after Folk and Ward [2]). (a) mean vs standard deviation, (b) mean vs skewness, (c) mean vs kurtosis, (d) skewness vs standard deviation, (e) standard deviation vs kurtosis and (f) skewness vs kurtosis.

#### 6.4. Skewness versus Standard Deviation

The plot between skewness and standard deviation produces a scattered trend in the form of nearly circular ring [2]. It happens due to either unimodal samples with good sorting or equal mixture of two different modes. However, the scattering of present sediments shows even clustering of grains in all sector (Figure 5d), that suggests the dominance of sand (coarse to very coarse) mode having subordinate silt. The presence of the silt is responsible for the skewness value to deviate into the negative proportion of the plot.

#### 6.5. Standard Deviation versus Kurtosis

Standard deviation versus kurtosis was plotted and correlated with standard plot proposed by [2], which is again followed by mixing of the varying proportions of two size-modes. Worst sorting is found in the bimodal mixtures with unequal amounts of two modes, and these also have lowest kurtosis [2]. In the present case the scatters very slightly differ from the pure sand region of the original curve of [2], indicating presence of small amount of fine-grained sediment. Most of the samples are leptokurtic along with platykurtic and mesokurtic and moderately sorted because of the dominance of sand (coarse to very coarse) size fraction (Figure 5e).

#### 6.6. Skewness versus Kurtosis

The plot between skewness versus kurtosis depends on two modes and follows a regular path as the mean-size changes [2]. In the present case all the values dominantly fall in an area represented by nearly pure sands, with less than 1% gravel and less than 5% silt of the establish plot of [2] (Figure 5f).

#### 7. C-M PLOT

In the present study C-M plot proposed by [30] is used to understand the dominant mode of transportation and the environment of deposition. All the hydrodynamic forces working during the deposition of the sediments have been analysed and interpreted by C-M plot [13, 30, 31]. Discussions on parameters C and M show that these parameters are indicators of hydraulic conditions under which sediments are deposited. In C-M diagrams, C is the one-percentile, M is the median of the grain-size distribution, characterize the coarsest fractions of the samples [31]. Percentile of the size distribution (C) in microns has been plotted against the median of the size distribution in microns (M) for the sediments (Figure 6). It is observed that all

the river sediments are above the CR line existing in segment N-O of the pattern. It indicates fairly well sorted sediments, almost entirely transported by rolling. Thus, the sediments dominated by coarse and very coarse sand deposited by the process of rolling in the Khurar River.

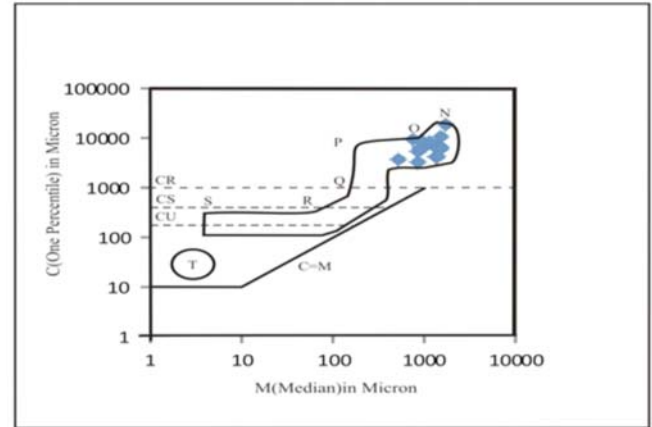
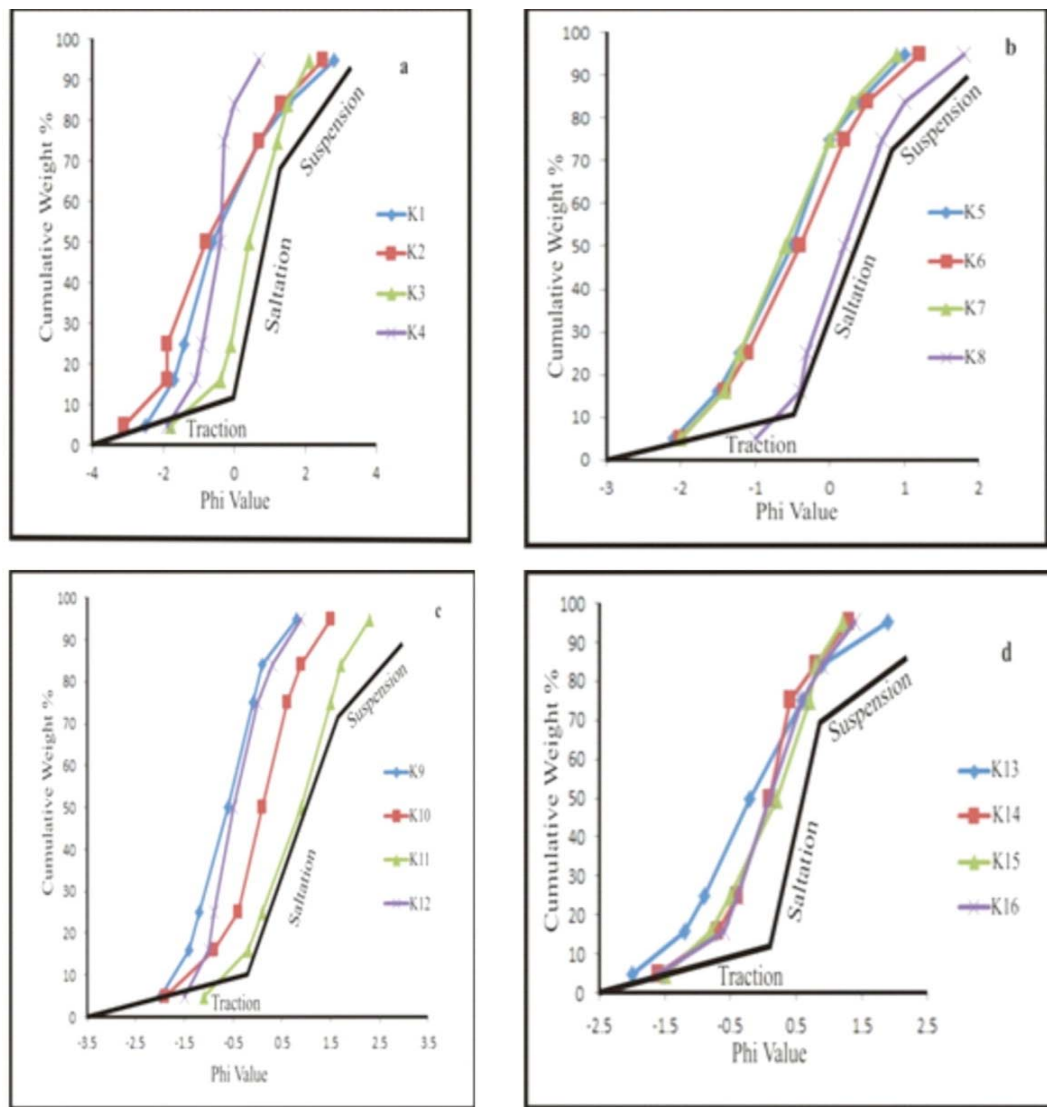


Figure 6: C-M plot showing concentration of sample points in upper left corner.

#### 8. LOG-PROBABILITY CURVES

Log-probability curves proposed by [27] had been plotted on the log probability (ordinate) paper in which Phi values were plotted on X axis and Y axis represents the cumulative frequency percentage of constituent grains. These plots indicate the mode of transportation of sediments within a depositional medium [27, 19]. This plot indicates two or three straight line rather than a single straight line. Each segment of the curve is interpreted to show different sub-population of grains that are transported simultaneously but by different ways i.e. suspension, saltation and traction bed load. Grain size probability plots for river sands are uniquely segmented into three differently sloped components because river move material in three ways: by suspension, by saltation and as bed-load [27]. All the sixteen samples have been plotted on log probability paper and compared with established trend for modern and ancient fluvial deposits proposed by [27], generally showing the dominance of saltation with traction and suspension domain also (Figure 7). This plot also indicates sorting of the sediment, where most of the sediments are moderately shorted probably due to prolonged transport and winnowing processes[29, 32]. In the present case, the sediments show moderate sorting even in the coarse of short distance may be because of winnowing and selective sorting of the sediments within the river.



**Figure 7:** Cumulative frequency curves (a to d) showing the traction, saltation and suspension populations.

## 9. CONCLUSIONS

Grain-size analysis of sediment samples from sixteen different sites representing coarse to very coarse-grained sand of Khurar River of Khajuraho area has been carried out. The important conclusions drawn are as follows:

- i) The frequency curves are dominantly indicative of coarse to very coarse-grained nature of the sediments.
- ii) The graphic mean value indicates more or less equal distribution of both very coarse and coarse sand-size particles.
- iii) In general, the samples show poor sorting in the beginning followed by moderate sorting to moderately well sorting at last site where the river confluences with river Ken. In most of the samples, both peak and tails are equally sorted giving rise to mesokurtic nature.
- iv) All the bivariate plots between mean, skewness, kurtosis and standard deviation are indicative of bimodal nature of sediments where sand-size (very coarse to coarse) dominates with subordinate silt.
- vi) All the sediments of Khurar River are characterized by the rolling process of deposition.
- vii) The log-probability curves of samples are showing dominance of saltation with traction and suspension as proposed by [27].
- viii) The change of sorting from poorly sorted to moderately sorted and moderately well sorted



suggests that the winnowing and selective sorting is possible even in the small river depending upon the hydrodynamic condition of the river.

## REFERENCES

- [1] Flemming BW. The influence of grain-size analysis methods and sediment mixing on curve shapes and textural parameters: Implications for sediment trend analysis. *Sedi Geol* 2007; 202: 425-35.  
<http://dx.doi.org/10.1016/j.sedgeo.2007.03.018>
- [2] Folk RL, Ward WC. Brazos River bar: a study in the significance of grain size parameters. *J Sedi Petrol* 1957; 27: 3-26.  
<http://dx.doi.org/10.1306/74D70646-2B21-11D7-8648000102C1865D>
- [3] Friedman GM. Differences in size distributions of populations of particles among sands of various origins. *Sedimentology* 1979; 26: 3-32.  
<http://dx.doi.org/10.1111/j.1365-3091.1979.tb00336.x>
- [4] Bui EN, Mazullo J, Wilding LP. Using quartz grain size and shape analysis to distinguish between aeolian and fluvial deposits in the Dallol Bosso of Niger (West Africa). *E Sur Proc Lan* 1990; 14: 157-66.  
<http://dx.doi.org/10.1002/esp.3290140206>
- [5] Zhanqiao W, Zhongyuan C, Maotian L, Jing C, Yiwen Z. Variations in downstream grain-sizes to interpret sediment transport in the middle-lower Yangtze River, China: A pre-study of Three-Gorges Dam. *Geomorphology* 2009; 113: 217-29.  
<http://dx.doi.org/10.1016/j.geomorph.2009.03.009>
- [6] Sarkar A, Paul DK, Potts PJ. Geochronology and geochemistry of mid-Archaean trondhjemitic gneiss from the Bundelkhand Craton, Central India. *Rec Res Geol* 1996; 16: 76-92.
- [7] Mondal MEA, Goswami JN, Deomurari MP, Sharma KK. Ionmicroprobe <sup>207</sup>Pb/<sup>206</sup>Pb ages zircons from Bundelkhand massif, north India: implications for crustal evolution of the Bundelkhand-Aravalli protocontinent. *Pre Res* 2002; 117: 85-100.  
[http://dx.doi.org/10.1016/S0301-9268\(02\)00078-5](http://dx.doi.org/10.1016/S0301-9268(02)00078-5)
- [8] Folk RL. *Petrology of Sedimentary Rocks*. Hemphill Austin, Texas 1980; 159.
- [9] Reineck HE and Singh IB. *Depositional Sedimentary Environments*, 2<sup>nd</sup> ed. Springer-Verlag-Berlin, Heidelberg, New York 1980; pp. 543.  
<http://dx.doi.org/10.1007/978-3-642-81498-3>
- [10] Pettijohn FJ. *Sedimentary Rocks*. 3<sup>rd</sup> ed. CBS Publishers, New Delhi 1984; pp. 628.
- [11] Lindholm RC. *A Practical Approach to Sedimentology*. George Allen & Unwin Publishers, London 1987; pp. 270.
- [12] Sengupta SM. *Introduction to Sedimentology*. Oxford & IBH Publishers, New Delhi, 1996; pp. 305.
- [13] Passega R. Texture as a characteristic of clastic deposition. *Am Asso Petro Geol* 1957; 41: 1952-84.
- [14] Udden JA. Mechanical composition of clastic sediments. *Geo Soc of Am Bull* 1914; 25: 655-744.
- [15] Wentworth CK. Method of computing mechanical composition type sediment. *Geo Soc Am Bull* 1929; 40: 711-90.  
<http://dx.doi.org/10.1130/GSAB-40-771>
- [16] Otto GH. Modified logarithmic for mechanical arithmetic probability graph for interpretation of mechanical analysis of sediment. *J Sedi Petrol* 1939; 9: 62-76.
- [17] Keller WD. Size distribution of sand in some dunes, beaches and sandstones. *Am Asso Petro and Geol Bull* 1949; 29: 215-21.
- [18] Inman DL, Chamberlain TK. Particle-size distribution in near shore sediments, In: *Hough Spec Pub* 1955; 3: 106-29.
- [19] Srivastava AK, Mankar RS. Grain Size Analysis and Depositional Pattern of Upper Gondwana Sediments (Early Cretaceous) of Salbardi Area, Districts Amravati, Maharashtra and Betul, Madhya Pradesh. *J Geol Soc India* 2008; 73: 393-406.  
<http://dx.doi.org/10.1007/s12594-009-0019-7>
- [20] Friedman GM. Distinction between dune, beach and river sands from their textural characteristics. *J Sedi Petrol* 1961; 31(4): 514-29.
- [21] Passega R. Texture as a characteristic of clastic deposition. *Am Asso Petro Geol* 1957; 41: 1952- 84.
- [22] Passega RS. Grain size representation by C-M pattern as a geological tool. *J Sedi Petrol* 1964; 34: 830-47.  
<http://dx.doi.org/10.1306/74D711A4-2B21-11D7-8648000102C1865D>
- [23] Friedman GM. On sorting, sorting coefficient and log-normality of the grain size distribution of sandstone. *J Geol* 1962; 70: 737- 57.
- [24] Friedman GM. Dynamic processes and statistical parameters compared for size frequency distribution of beach and river sands. *J Sedi Petrol* 1967; 37 (2): 327-54.
- [25] Sahu BK. Significance of grain size-distribution statistics in interpretation of depositional environment. *Pun Uni Res Bull* 1964; 15: 213-19.
- [26] Moiola RJ, Weiser D. Textural parameters: An evaluation. *J Sedi Petrol* 1968; 38 (1): 45-53.
- [27] Visher GS. Grain size distributions and depositional processes. *J Sedi Petrol* 1969; 39: 1074-1106.
- [28] Kalicki T. Grain size analysis of overbank deposits as carriers of paleogeographical information. *Quat In* 2000; 72: 107-114.  
[http://dx.doi.org/10.1016/S1040-6182\(00\)00026-4](http://dx.doi.org/10.1016/S1040-6182(00)00026-4)
- [29] Bartholdy J, Christiansen C, Pedersen JBT. Comparing special grain size trends inferred from textural parameters using percentile statistical parameters and those based on log- hyperbolic method. *Sedi Geol* 2007; 202: 436-52.  
<http://dx.doi.org/10.1016/j.sedgeo.2007.03.008>
- [30] Passega R, Byramjee R. Grain size image of clastic deposits. *Sedi Geol* 1969; 13: 180-90.
- [31] Passega R. Significance of C-M diagrams of sediments deposited by suspension. *Sedi Geol* 1977; 24: 723-33.  
<http://dx.doi.org/10.1111/j.1365-3091.1977.tb00267.x>
- [32] Bartholoma A, Flemming BW. Progressive grain-size sorting along an intertidal energy gradient. *Sedi Geol* 2007; 202: 464-472.  
<http://dx.doi.org/10.1016/j.sedgeo.2007.03.010>